

羌塘地块西部晚三叠世灰岩古地磁研究及其构造意义

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内容提要:青藏高原羌塘地块和拉萨地块汇聚—碰撞拼合过程的研究对认识青藏高原中部隆升历史及其动力学过程具有重要的科学意义, 而羌塘地块古地磁研究对理解上述问题至关重要。通过对羌塘地块西部改则地区晚三叠世灰岩的系统古地磁测定, 获得其高温剩磁分量。但是这一高温剩磁分量未通过褶皱检验, 表明为后期重磁化的结果。研究剖面高温特征剩磁平均方向在地理坐标下为 $Dg=349.3^\circ$, $Ig=40.4^\circ$, $\kappa g=45.4$, $\alpha_{95}=6.5^\circ$, 相应的古地磁极为 $76.4^\circ N, 311.1^\circ E$, $dp/dm=4.7^\circ/7.9^\circ$ 。这一古地磁极与羌塘地块早白垩世约 110~100 Ma 的古地磁极在古地磁误差范围内重合, 表明其重磁化的时代为早白垩世约 110~100 Ma。综合分析羌塘地块和拉萨地块古地磁结果, 并结合海相地层、蛇绿岩和洋岛等地质证据, 显示班公湖—怒江特提斯洋西段闭合的时间发生在早白垩世晚期约 110~100 Ma。改则地区晚三叠世灰岩的早白垩世晚期重磁化作用与羌塘/拉萨地块西部的碰撞密切相关。

关键词: 羌塘地块; 古地磁; 重磁化; 羌塘—拉萨地块碰撞; 班公湖—怒江洋

青藏高原经历了特提斯体系的长期演化, 是地质历史过程中高原上诸地块在不同时期碰撞—拼合至欧亚大陆南缘而形成的 (Sengör, 1987; Dewey et al., 1988; Yin An et al., 2000; Metcalfe, 2006, 2013; Xu Zhiqin et al., 2011; Wang Chengshan et al., 2014; Zhu Dicheng et al., 2016)。青藏高原诸地块间如何碰撞—拼合及其持续效应是研究青藏高原形成和演化的重要科学问题 (Yin et al., 2000; Metcalfe, 2006, 2013; Xu Zhiqin et al., 2011; Zhu Dicheng et al., 2013; Yan Maodu et al., 2016)。羌塘地块和拉萨地块的碰撞造成了大量的陆内构造缩短, 导致青藏高原中部在印度与亚洲大陆碰撞前就可能已有一定的海拔高度 (Murphy et al., 1997; Kapp et al., 2005, 2007; Volkmer et al., 2007, 2014), 形成白垩纪时期的青藏高原雏形 (Wang Chengshan et al., 2014; Xu Zhiqin et al., 2016)。另外, 班公湖—怒江缝合带不仅是一条重要

的构造界线, 而且是一条重要的成矿带, 具有巨大的成矿潜力 (Song Yang et al., 2014; Gao Ke et al., 2017; Liu Hong et al., 2017; Tang Juxing et al., 2017; Zheng Youye et al., 2017)。班公湖—怒江缝合带所代表的洋盆的闭合时间是重要的成矿期。因此, 羌塘地块和拉萨地块汇聚—碰撞拼合过程 (班公湖—怒江特提斯洋演化和闭合过程) 的研究不仅对认识青藏高原中部早期隆升历史和动力学过程具有重要的科学意义, 而且还可为查明班公湖—怒江成矿带区域成矿规律提供清楚明确的地质构造背景。

前人根据不同研究方法对羌塘地块和拉萨地块汇聚—碰撞过程和班公湖—怒江洋演化及闭合过程进行了深入研究和探讨, 并取得了丰富的研究成果, 但是至今还存在较大争议。例如, 不同学者提出的羌塘地块与拉萨地块碰撞及班公湖—怒江特提斯洋闭合的时代自中—晚侏罗世 (Xu Ronghua et al.,

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1985; Dewey et al. , 1988; Li Shun et al. , 2017; Ma Anlin et al. , 2017, 2018; Sun Gaoyuan et al. , 2019)、早白垩世(Kapp et al. , 2005, 2007; Mo Xuanxue et al. , 2006; Zhu Dicheng et al. , 2016; Huang Tongtong et al. , 2017; Ma Yiming et al. , 2018; Lei Chuanyang et al. , 2019)至晚白垩世(Zhang Kaijun et al. , 2012; Li Yalin et al. , 2013, 2015; Liu Deliang et al. , 2014; Liu Weiliang et al. , 2014; Fan Jianjun et al. , 2014, 2015; Wang Baodi et al. , 2016; Chen Weiyan et al. , 2018)不等。羌塘地块与拉萨地块碰撞为近南北向的陆-陆碰撞,那么定量约束两块体碰撞时和碰撞前古纬度位置对认识羌塘地块与拉萨地块汇聚-碰撞过程和班公湖-怒江特提斯洋的演化及闭合过程等科学问题有重要意义。古地磁学能定量地约束板块的古纬度,在地块古地理位置重建和地块运动学的研究上具有独特的优势,是制约板块间相互作用最理想方法之一。但是,目前羌塘地块古地磁学研究相对比较薄弱,尤其是中生代可靠的古地磁数据非常有限。另外,现有可靠的古地磁结果主要来自羌塘地块中一东部地区(Cheng Xin et al. , 2012; Song Chunyan et al. , 2012; Ren Haidong et al. , 2013;

Song Peiping et al. , 2015, 2017; Tong Yabo et al. , 2015; Yan Maodu et al. , 2016; Ran Bo et al. , 2017; Meng Jun et al. , 2018; Cao Yong et al. , 2019; Ma Yiming et al. , 2019)(图 1),而由于印度与亚洲大陆碰撞以及青藏高原中部大量走滑断裂作用的影响,白垩纪以来青藏高原中部不同地区之间都存在不同的局部水平旋转作用,就导致羌塘地块中一东部地区的古地磁极数据很难用来约束羌塘地块西部的古地理位置及其运动过程。本文选择羌塘地块西部改则地区晚三叠世灰岩开展古地磁研究,以期为羌塘地块与拉萨地块汇聚-碰撞过程和班公湖-怒江特提斯洋演化及闭合过程提供古地磁定量约束。

1 地质背景及样品采集

青藏高原南起喜马拉雅山脉南缘,北至阿尔金和祁连山北缘,西临帕米尔高原和喀喇昆仑山脉,东接四川盆地。青藏高原可根据多条缝合带划分成多个不同块体(图 1),从南至北依次为喜马拉雅地块、拉萨地块、羌塘地块、松潘-甘孜-可可西里地块、昆仑地块-柴达木地块和祁连地块。羌塘地块位于青藏高原腹地,东西长约 1200 km,中部宽约 500~

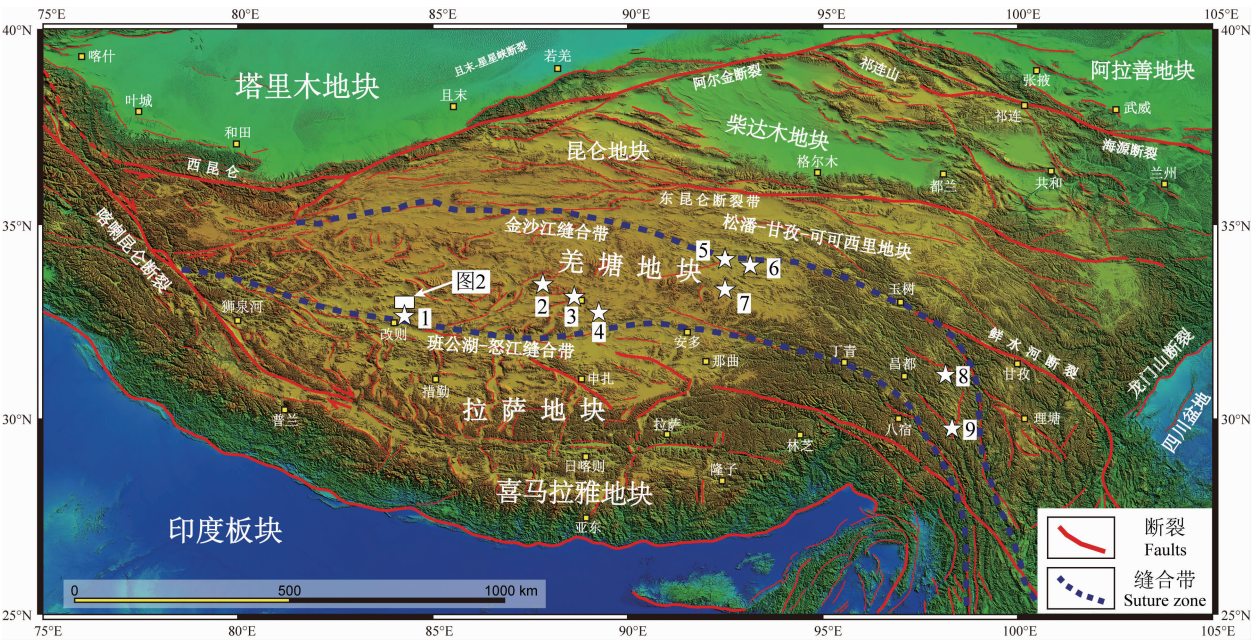


图 1 青藏高原及其周围构造简图

Fig. 1 Schematic tectonic map of the Tibetan Plateau and the adjacent region

图中五角星代表现有羌塘地块古地磁结果研究位置分布(the star represent the paleomagnetic research location in Fig. 1):1—Chen Weiwei et al. , 2017; 2—Song Chunyan et al. , 2012; 3—Meng Jun et al. , 2018; 4—Cao Yong et al. , 2019; 5—Song Peiping et al. , 2015, 2017; Ma Yiming et al. , 2019; 6—Halim et al. , 1998; 7—Cheng Xin et al. , 2012; Ren Haidong et al. , 2013; Yan Maodu et al. , 2016; Ran Bo et al. , 2017; 8—Tong Yabo et al. , 2017; 9—Huang Kainian et al. , 1992; Tong Yobo et al. , 2015

600 km,西部和东部宽度小于 150 km(Yin et al., 2000)。羌塘地块南以班公湖—怒江缝合带为界与拉萨地块相邻,北以金沙江缝合带为界与松潘—甘孜—可可西里地块相邻(图 1)。近年来,地质学家们在羌塘地块中西部变质杂岩带中,沿龙木错—双湖一线发现古特提斯阶段的蛇绿混杂岩、榴辉岩和蓝片岩的各种组分,且发现该带两侧的古生物面貌和沉积地层完全不同(Xu Zhiqin et al., 2013)。Li Cai(1987)曾提出该带为一条北西西—南东东向的板块缝合带(龙木错—双湖—澜沧江板块缝合带),并提出该缝合带为冈瓦纳大陆的北界的观点。许多学者根据龙木错—双湖—澜沧江板块缝合带的发现将羌塘地块划分为南羌塘地块和北羌塘地块(Li Cai et al., 2008, 2009; Peng Touping et al., 2015),也有学者称之为西羌塘和东羌塘(Zhang Kaijun et al., 2012; Zhu Dicheng et al., 2013, 2016)。

本次研究区位于羌塘地块西部改则县城东北约 90 km 处玛日玛地区,在班公湖—怒江缝合带北侧,龙木错—双湖缝合带以南,构造上属于南羌塘地块(图 1)。区内出露的地层主要是三叠系地层,侏罗系、古近系和新近系地层零星出露(图 2a)。上三叠统日干配错组(T_3r)地层岩石组合以微晶灰岩和鲕粒灰岩为主,产双壳类和珊瑚类等化石(Geological Survey of Tibet Autonomous Region, 2006)。下一中侏罗统色哇组(J_{1-2s})地层岩石总体以泥质岩为主,产珊瑚化石(Geological Survey of Tibet Autonomous Region, 2006),区内与下伏日干配错组地层呈断层接触关系。纳丁错组地层岩性以中—基性火山岩为主,区内与日干配错组以角度不整合接触,研究显示研究区内纳丁错组火山岩的时代为 ~ 110 Ma(Kapp et al., 2005; Chang Qingsong et al., 2011)。新近系康托组(Nk)岩石以砾岩夹砂岩为主,区内与日干配错组和色哇组以角度不整合接触。区内东部出露燕山晚期花岗闪长岩($\gamma\delta K_1$)。Geological Survey of Tibet Autonomous Region (2006)曾在研究区内对日干配错组地层进行了实测地层剖面,实测剖面显示日干配错组岩性以微晶灰岩和鲕粒灰岩为主,其间不同程度夹有角砾状灰岩、砂屑灰岩、泥灰岩、粒屑灰岩、藻灰岩、介屑灰岩。在改则其他区域的日干配错组地层中发现生物礁灰岩为海绵礁和珊瑚礁,鉴定结果为晚三叠世(Geological Survey of Tibet Autonomous Region, 2006)。

本次古地磁采样工作在玛日玛村至扎美仁村公路沿线一带日干配错组地层内开展。采样剖面地层

出露较好,岩石新鲜,地层产状清晰(图 2b~2e),共采集了 23 个古地磁采点约 224 块岩心样品。剖面北部采点 MT1~MT3 所在地层朝 SSW 倾,采点 MT4~MT10 所在地层朝 NNE 倾,采点 MT13~MT20 所在地层朝 SSE 倾,采样剖面地层总体上地层倾向方向变化较大。另外,采样剖面地层倾角也变化较大,倾角从 $\sim 40^\circ$ 至近直立。

2 古地磁测试与分析

全部样品在室内加工成标准古地磁样品,并利用 TD-48 大型热退磁炉进行系统热退磁处理。热退磁温度在低温段间隔为 $50\sim 60^\circ\text{C}$,高温段间隔为 $20\sim 30^\circ\text{C}$ 。样品的剩磁测量利用 2G-755R 超导磁力仪进行。为确定岩石中样品的主要载磁矿物,对代表样品进行了三轴等温剩磁热退磁实验,测试仪器为 JR-6A 旋转磁力仪。以上测试均在自然资源部古地磁与古构造重建重点实验室磁屏蔽屋内进行。磁屏蔽屋内的磁场 < 300 nT,排除了外界磁场对样品测试数据的影响。样品的退磁结果分析利用主向量法(Kirschvink, 1980),最后以采样点为单位进行 Fisher 统计分析(Fisher, 1953)。

2.1 三轴等温剩磁热退磁结果

对日干配错组灰岩代表样品的 Z 轴、Y 轴和 X 轴方向依次施加 2.5 T、0.4 T 和 0.12 T 磁场,然后对样品进行系统热退磁实验。三轴等温剩磁热退磁实验结果显示,样品 MT3-2 和 MT11-5A 的三轴磁性成分的剩磁强度均在 $\sim 680^\circ\text{C}$ 时衰减至接近为零(图 3a, 3c);样品 MT8-6A 和 MT20-4 的硬磁成分的剩磁强度在 $\sim 680^\circ\text{C}$ 时衰减至接近为零,而中间磁成分和软磁成分的剩磁强度在 $\sim 580^\circ\text{C}$ 时接近为零(图 3b, 3d)。以上结果表明日干配错组样品中的主要载磁矿物为赤铁矿和磁铁矿。

2.2 系统热退磁结果

日干配错组灰岩样品系统热退磁结果显示大多数样品只存在单一剩磁分量(图 4a, 4d, 4g, 4h)。样品在 180°C 至 580°C 或 680°C 温度段之间可分离出一组高温剩磁分量。日干配错组灰岩古地磁样品数据经分析后,共获得了 12 个采点的高温特征剩磁分量(表 1)。12 个采样点高温特征剩磁分量的平均磁化方向为: $D_g = 349.3^\circ$, $I_g = 40.4^\circ$, $\kappa_g = 45.4$, $\alpha_{95} = 6.5^\circ$ (地理坐标下); $D_s = 247.0^\circ$, $I_s = 83.6^\circ$, $\kappa_s = 2.2$, $\alpha_{95} = 39.0^\circ$ (层面坐标下)(表 1, 图 5)。

日干配错组样品高温特征剩磁分量经产状校正后变得更加分散($k_g/k_s = 20.6$)(图 5)。另外,古地

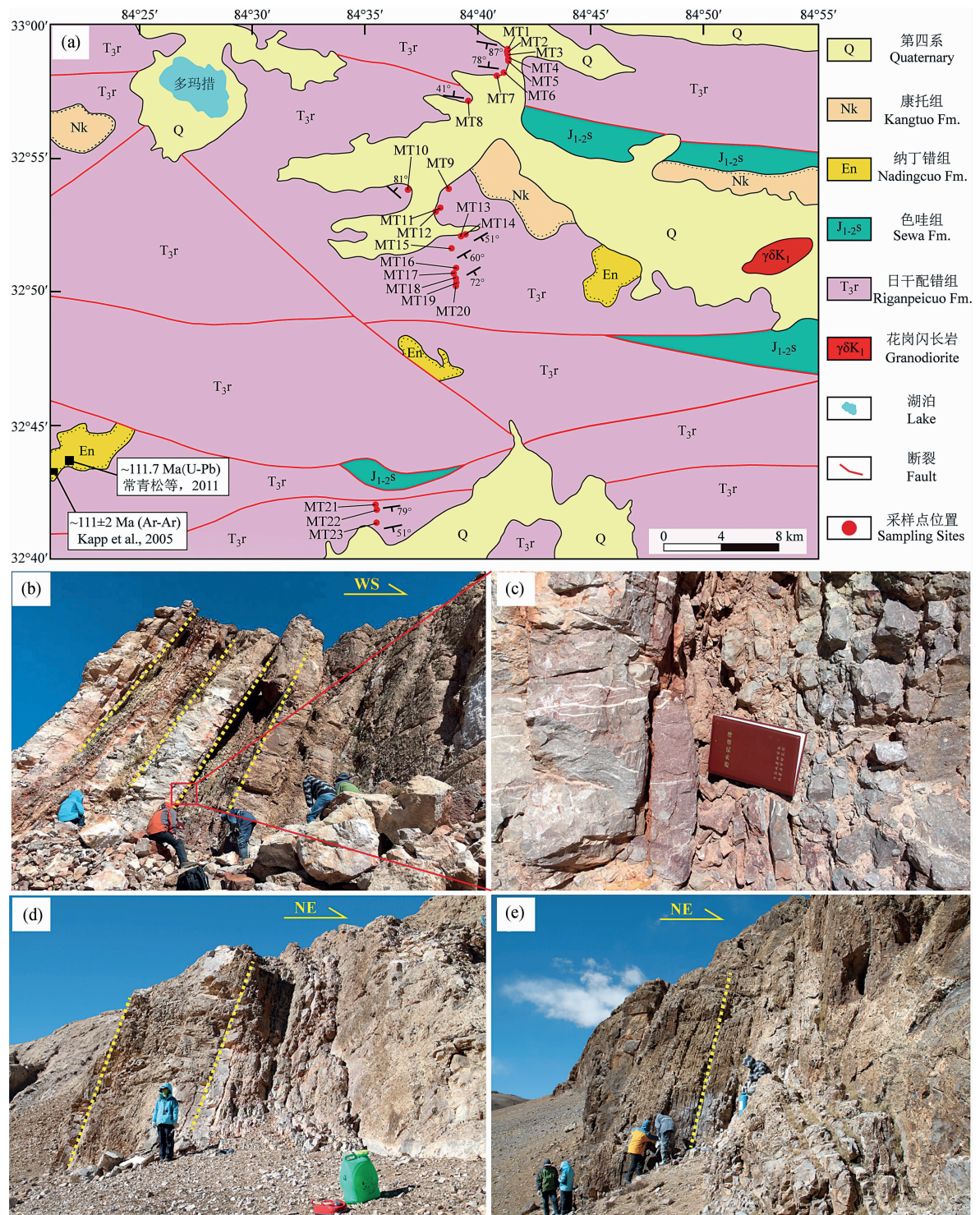


图 2 羌塘地块西部玛日玛地区地质简图(a)和上三叠统日干配错组地层野外露头及采样照片(b~e)

Fig. 2 Simplified geological map of the Marima area in the western portion of the Qiangtang block (a) and photographs showing field outcrops and sampling sites of the Upper Triassic Riganpeicuo Formation (b~e)

磁数据未通过褶皱检验($\xi_{in-situ} = 2.884 < \xi_c = 4.036 < \xi_{tilt-corrected} = 8.47$ 在 95% 置信区间为负褶皱检验;
 $\xi_{in-situ} = 5.308 < \xi_c = 5.624 < \xi_{tilt-corrected} = 11.64$ 在 99% 置信区间也为负褶皱检验)(McFadden, 1990)。高温特征剩磁分量的逐步褶皱展平检验

(Watson et al., 1993)结果显示 Fisher 统计精度参数 k 的极大值出现在 -1.8% 展平状态,表明日干配错地层单元的特征剩磁是在褶皱变形之后获得的。以上结果显示本次研究从日干配错组地层中获得的晚三叠世古地磁数据为后期重磁化结果。

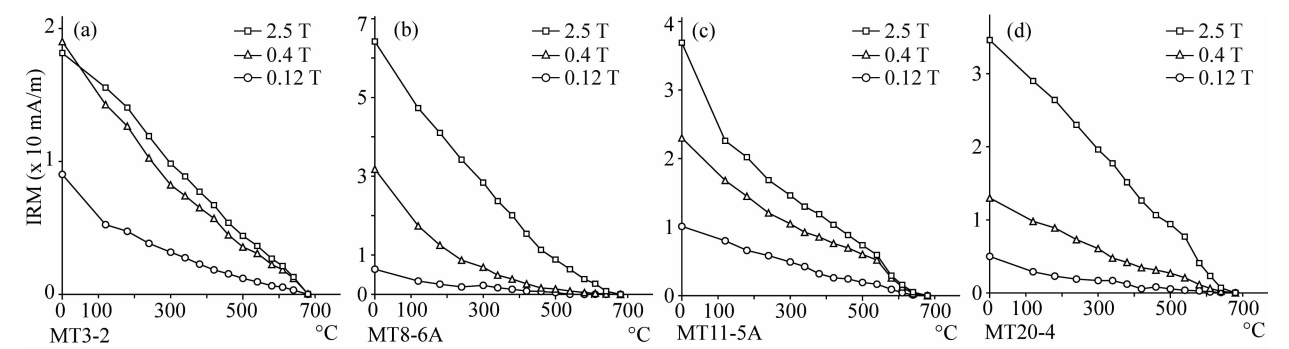


图 3 羌塘地块西部上三叠统日干配错组岩石代表样品三轴等温剩磁热退磁结果图

Fig. 3 Thermal demagnetization of three-component isothermal remanent magnetization for representative samples from the Upper Triassic Riganpeicuo Formation in the western portion of the Qiangtang block

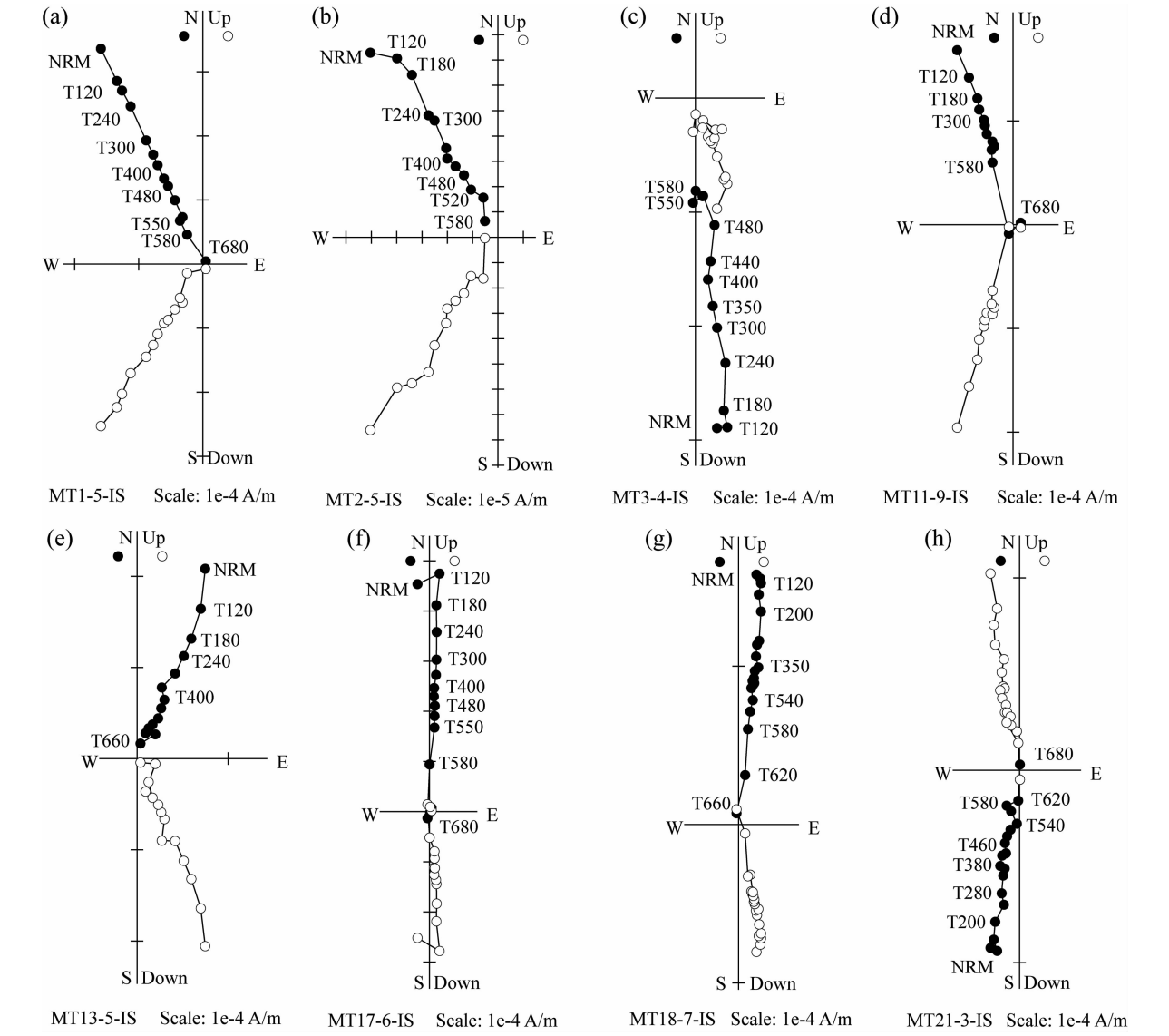


图 4 羌塘地块西部上三叠统日干配错组岩石代表样品 Z 矢量图(地理坐标下)

Fig. 4 Orthogonal vector diagrams of representative samples from the Upper Triassic Riganpeicuo Formation in the western portion of the Qiangtang block (in situ coordinates)

实心圆圈和空心圆圈分别代表剩磁方向在水平面和铅垂面上的投影

The solid and open circles represent vectors endpoints projected onto horizontal and vertical planes, respectively

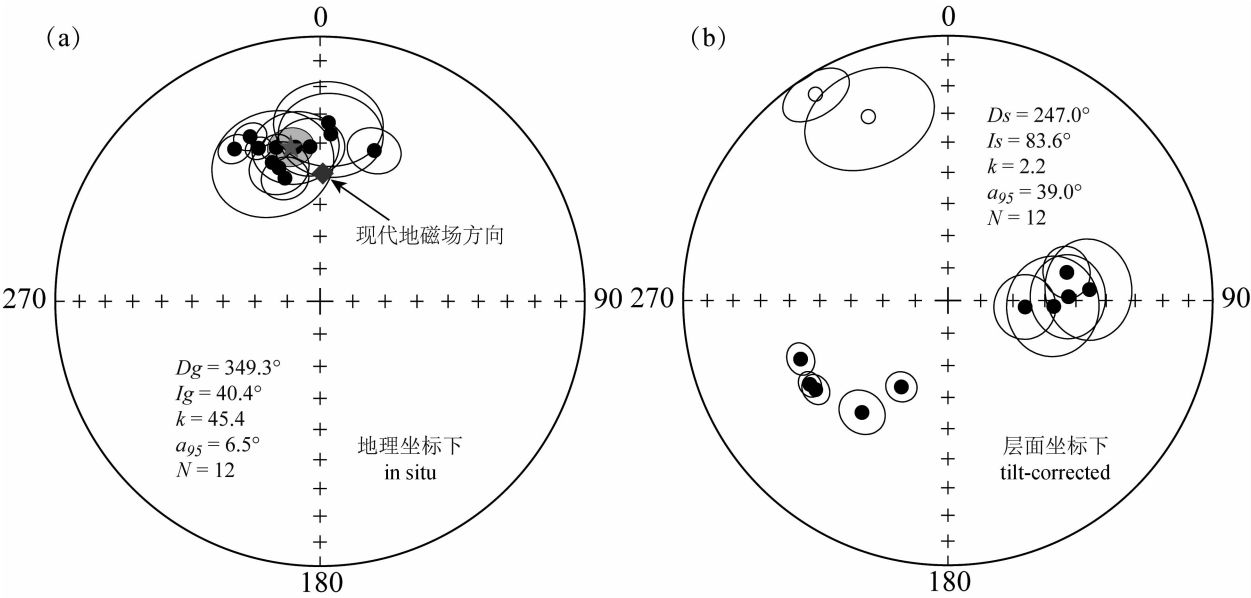


图 5 羌塘地块西部上三叠统日干配错组岩石样品高温特征剩磁分量地理坐标下(a)和层面坐标下(b)赤平投影图
Fig. 5 In situ (a) and tilt-corrected (b) equal-area projections of site-mean directions from the Upper Triassic Riganpeicuo Formation in the western portion of the Qiangtang block

表 1 羌塘地块西部上三叠统日干配错组岩石样品
高温特征剩磁分量结果统计表
Table 1 Site-mean paleomagnetic directions of the high-temperature components from the Upper Triassic Riganpeicuo Formation in the western portion of the Qiangtang block

采点	走向/ 倾向	<i>n</i> / <i>N</i>	<i>Dg</i> (°)	<i>Ig</i> (°)	<i>Ds</i> (°)	<i>Is</i> (°)	<i>κ</i>	<i>α</i> ₉₅ (°)
MT1	110/87	7/11	337.0	32.9	248.3	39.9	172.2	4.6
MT2	108/88	9/10	338.1	37.7	238.7	38.9	190.3	3.7
MT3	110/89	7/8	344.1	39.4	236.0	39.6	190.8	4.4
MT9	205/65	4/6	342.9	46.2	327.3	-8.2	114.0	8.6
MT11	230/72	5/9	341.0	43.8	336.6	-25.3	20.1	17.5
MT13	60/51	10/10	19.8	39.3	76.7	51.9	41.8	7.6
MT15	60/60	6/10	356.3	41.1	94.8	66.1	48.6	9.7
MT17	49/74	8/10	3.7	36.8	85.6	45.4	15.8	14.4
MT18	45/71	9/10	350.9	40.7	88.2	52.4	18.8	12.2
MT19	60/72	6/8	2.7	32.9	93.2	57.1	21.4	14.8
MT21	82/79	6/9	330.7	34.4	208.3	59.6	201.4	4.7
MT22	102/78	4/9	344.1	50.0	217.5	45.7	177.5	6.9
Mean		12(81)	349.3	40.4	—	—	45.4	6.5
			—	—	247.0	83.6	2.2	39.0

注：*n*/*N* 为参加统计/剩磁测量样品数；*Dg*, *Ig* (*Ds*, *Is*) 分别为地理坐标下(层面坐标下)特征剩磁的偏角和倾角；*κ*, *α*₉₅ 分别为 Fisher 统计精度参数和平均方向的 95% 置信圆锥半顶角。

3 讨论

3.1 重磁化时代

Van der Voo (1990) 提出了 7 条古地磁数据可靠性判别标准, 我们在此基础上重点考察羌塘地块

白垩纪以来的古地磁结果是否通过稳定性检验和是否经过构造校正, 从而筛选出可靠地古地磁数据(表 2)。值得注意的是, Chen Weiwei et al. (2017) 发表的早白垩世 110~100 Ma 的古地磁结果中 GZ35-36、GZ37、GZ38 和 GZ39-41 四个点的古地磁结果明显偏离其他 10 个采点的结果(见 Chen Weiwei et al., 2017 图 8f), 因此这 4 个采点的结果不再进行分析。其余 10 个采点的古地磁结果的平均方向为 *Ds*/*Is*=346.9°/42.0°, *a*₉₅=5.1°; 相应的古地磁极位置为 75.8°N/321.9°E, *dp/dm*=3.8°/6.3°。这一古地磁结果也通过褶皱检验($\xi_{in-situ}$ =6.327> ξ_c =3.685> $\xi_{tilt-corrected}$ =3.492 在 95% 置信区间为正褶皱检验; $\xi_{in-situ}$ =6.327> ξ_c =5.12> $\xi_{tilt-corrected}$ =3.492 在 99% 置信区间也为正褶皱检验)。此外, Meng Jun et al. (2018) 对羌塘地块双湖县城附近上白垩统阿布山组砂岩开展了古地磁学研究, 从 70 余块样品中获得了相应的中温分量和高温分量。Meng Jun et al. (2018) 将中温分量和高温分量分别进行 *E*/*I* 矫正分析。中温分量平均磁倾角为 46.1°±2.9°, 中温分量 *E*/*I* 矫正结果平均倾角为 48.8°(置信范围为 45.1°~59.6°), 表明其中温分量结果未受到明显的磁倾角浅化作用的影响。高温分量平均磁倾角为 37.2°±3.7°, 高温分量 *E*/*I* 矫正结果平均倾角为 42.8°(置信范围为 36.3°~51.2°)。Meng Jun et al. (2018) 认为其研究结果高温分量磁倾角发生

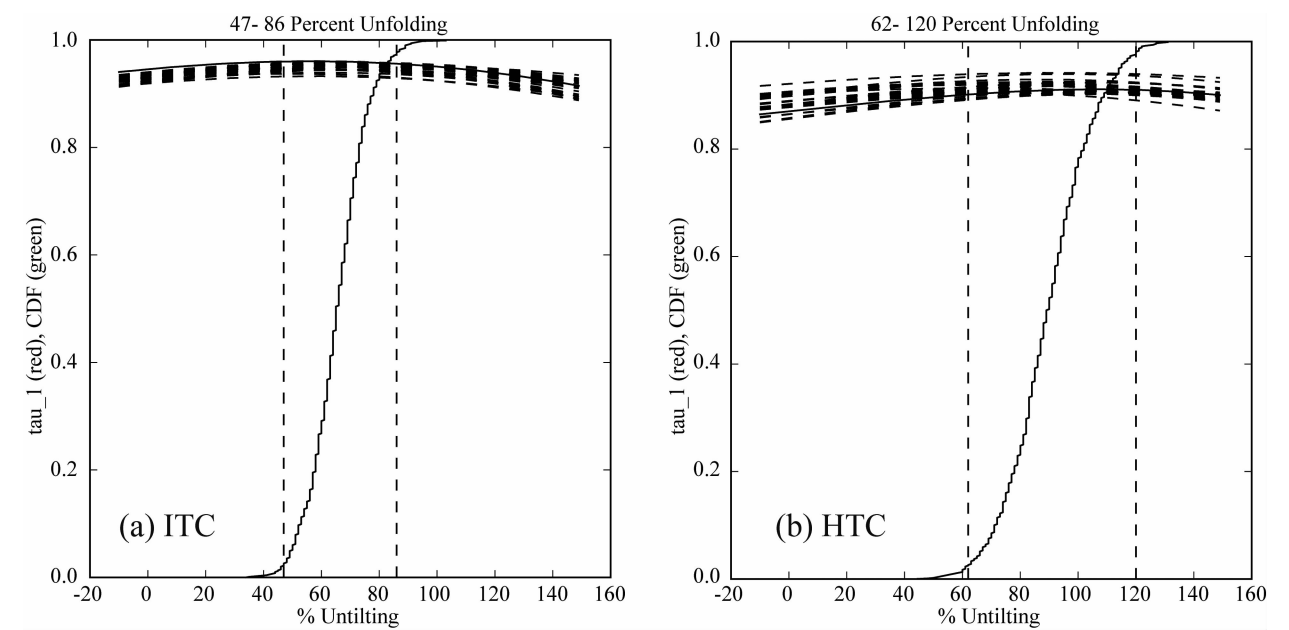


图 6 Meng Jun et al. (2018) 研究中温分量(a)和高温分量(b)褶皱检验结果图

Fig. 6 Nonparametric fold tests of the Intermediate Temperature Component (ITC) and High Temperature Component (HTC) directions from Meng Jun et al. (2018)

了浅化现象,因此其用中温分量结果来限定羌塘地块晚白垩世的古纬度。但是其高温分量的原始平均磁倾角($I=37.2^{\circ}\pm3.7^{\circ}$)却在 E/I 矫正结果的磁倾角置信范围内,并不能说明其研究地层磁倾角发生了明显浅化现象。另外,我们对 Meng Jun et al. (2018)所发表的结果进行基于样品数据的褶皱检验(Tauxe et al., 1994)。褶皱检验结果显示其中温分量很可能在褶皱形成过程中获得的(图 6a),而其高温分量则可能是在褶皱前获得的(图 6b)。因此,我们参考 Meng Jun et al. (2018)所报道的羌塘地块

晚白垩世古地磁研究的高温特征剩磁分量数据结果来进行进一步讨论。

本次研究从羌塘地块西部改则地区上三叠统日干配组灰岩中获得磁化方向为: $Dg=349.3^{\circ}$, $Ig=40.4^{\circ}$, $\kappa g=45.4$, $\alpha_{95}=6.5^{\circ}$,相应的古地磁极为 76.4°N , 311.1°E , $dp/dm=4.7^{\circ}/7.9^{\circ}$ 。与羌塘地块白垩纪以来可靠的古地磁极数据对比,显示本次研究获得的古地磁极与羌塘地块晚白垩世古地磁极(Huang Kainian et al., 1992; Tong Yabo et al., 2015; Chen Weiwei et al., 2017; Meng Jun et al.,

表 2 羌塘地块白垩纪—新近纪古地磁极统计表

Table 2 Summary of the Cretaceous-Neogene paleomagnetic results from the Qiangtang block

SLat. (°N)	SLon. (°E)	时代 (Ma)	N/n	Plat. (°N)	Plon. (°E)	$A_{95}(dp/dm)$ (°)	检验	古纬度 (°N)	参考文献
地块西部									
34.5	92.7	E	7/31	72.2	242.1	7.7	F,R	15.8 ± 9.4	Halim et al., 1998
33.2	88.7	K ₂	9/71	29.2	171.8	2.5/4.3	F	16.9 ± 3.9	Meng Jun et al., 2018
32.5	84.3	K ₂	22/174	45.4	348.1	3.1	F	18.8 ± 3.6	Chen Weiwei et al., 2017
32.5	84.3	110~100	10/57	75.8	321.9	3.8/6.3	F	24.3 ± 5.1	ChenWeiwei et al., 2017
32.9	84.7	110~100	12/81	76.4	311.1	4.7/7.9	—	22.8 ± 6.5	本次研究
32.9	83.5	120~115	16/127	37.9	162.2	5.1	—	27.8 ± 4.9	Cao Yong, 2018
地块东部									
31.0	98.2	E ₂	43/352	57.9	192.1	2.9	F,R	18.4 ± 3.4	Tong Yabo et al., 2017
29.7	98.4	K ₂	17/186	47.0	165.1	6.7/9.3	F,R	28.8 ± 7.5	Tong Yabo et al., 2015
29.7	98.7	K ₂	11/64	56.7	172.7	10.6	F	27.4 ± 10.4	Huang Kainian et al., 1992
29.7	98.4	K ₁	12/56	40.6	170.5	13.0	F	22.9 ± 13.9	Huang Kainian et al., 1992

注:SLat. 和 SLon. 分别为采样点的纬度和经度;N/n 为采点数/样品数;Plat. 和 Plon. 分别为古地磁极的经度和纬度;A₉₅ 为为古地磁极的 95%置信圆锥半顶角;dp/dm 为 95% 置信椭圆半长轴与半短轴;F 代表通过褶皱检验;R 代表通过倒转检验;古纬度是以改则为参考点(32.5°N / 84°E)计算。

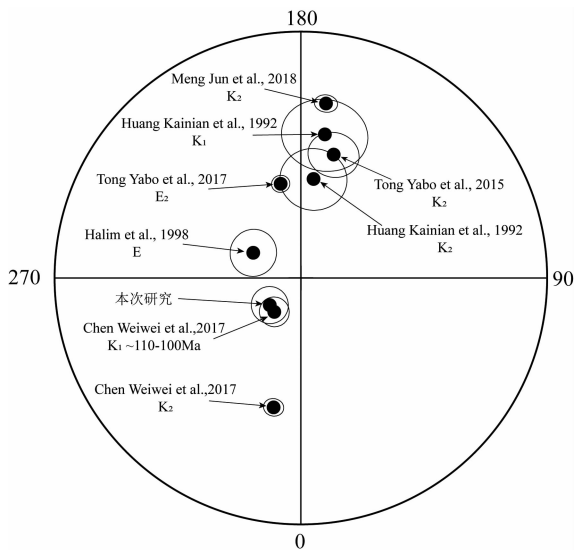


图 7 羌塘地块白垩纪—新近纪古地磁极赤平投影图
Fig. 7 Equal-area projection of the Cretaceous-Neogene paleomagnetic poles of the Qiangtang block

2018)和古近纪—新近纪古地磁极(Halim et al., 1998; Tong Yabo et al., 2017)均存在较大差异(图 7)。另外,本次研究从日干配错组灰岩获得的高温剩磁方向(地理坐标下: $D/I=349.3^{\circ}/40.4^{\circ}$)也与参考点现代地磁场的方向($D/I=1.3^{\circ}/50.2^{\circ}$)存在差异(图 5)。而本次研究获得的古地磁极与 Chen Weiwei et al. (2017)从羌塘地块改则地区早白垩世(110~100 Ma)火山岩中获得的古地磁极重叠(图 7)。以上分析表明羌塘地块西部改则地区晚三

叠世灰岩的重磁化时代应为早白垩世 110~100 Ma。

3.2 构造意义

本次研究从晚三叠世灰岩获得的重磁化结果与 Chen Weiwei et al. (2017)从早白垩世(110~100 Ma)火山岩中获得的古地磁结果一致,综合古地磁结果表明羌塘地块西部南缘在约 110~100 Ma 期间位于 $23.3 \pm 4.0^{\circ}\text{N}$ 。拉萨地块西部白垩纪古地磁研究结果见表 3,利用小圆弧拟合法对 Chen Weiwei et al. (2012)、Ma Yiming et al. (2014)和 Bian Weiwei et al. (2017)从拉萨地块西部早白垩世火山岩中获得的高质量古地磁极以改则为参考点($32.5^{\circ}\text{N} / 84^{\circ}\text{E}$)进行小圆弧拟合,获得拉萨地块西部北缘早白垩世的古纬度为 $20.1 \pm 2.4^{\circ}\text{N}$ (图 8a)。对比羌塘地块和拉萨地块西部早白垩世古纬度显示两地块西部地区在早白垩世晚期 110~100 Ma 的古纬度差为 $3.2^{\circ} \pm 3.1^{\circ}$,显示在古地磁误差范围内无明显差别。另外,拉萨地块西部晚白垩世古地磁极分布在以改则参考点为圆心以 $74.2^{\circ} \pm 2.7^{\circ}$ 为半径的圆弧上,对比羌塘地块晚白垩世古地磁结果显示两地块在晚白垩世时期古纬度已经重叠(图 8b)。而根据改则西北部早白垩世火山岩中获得的古地磁结果显示羌塘地块西部南缘在 120~115 Ma 时期的古纬度为 $27.8 \pm 4.9^{\circ}\text{N}$ (Cao Yong, 2018),表明在 120~115 Ma 羌塘地块和拉萨地块西部之间还存在 $7.2^{\circ} \pm 5.5^{\circ}$ 的古纬度差。

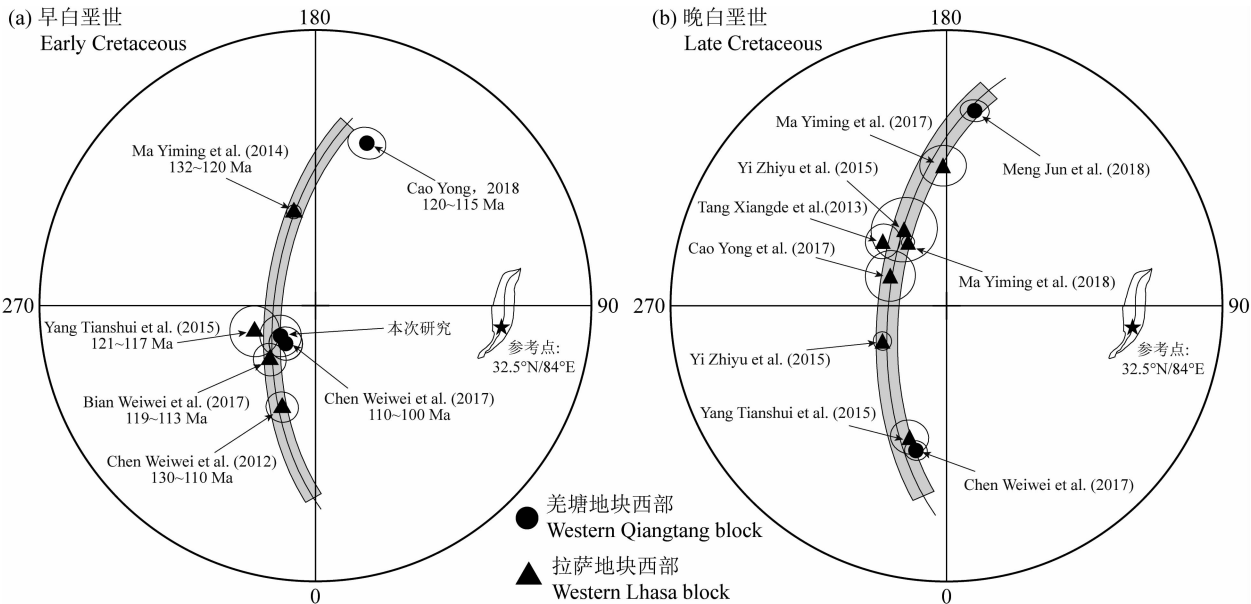


图 8 羌塘地块和拉萨地块西部早白垩世(a)和晚白垩世(b)古地磁极赤平投影图

Fig. 8 Equal-area projection showing the available Early (a) and Late (b) Cretaceous paleomagnetic poles obtained from the western portion of the Qiangtang and Lhasa blocks

表 3 拉萨地块西部白垩纪古地磁极统计表

Table 3 Summary of the Cretaceous paleomagnetic poles from the western portion of the Lhasa block

SLat. (°N)	SLon. (°E)	时代(Ma)	N/n	Plat. (°N)	Plon. (°E)	$A_{95}(dp/dm)^{(\circ)}$	古纬度(°N)	参考文献
32.6	80.2	~69	44/306	47.8	181.4	6.4	19.0 ± 7.4	Ma Yiming et al. , 2017
32.4	80.1	~92	10/78	64.1	209.0	9.6	15.8 ± 11.7	Yi Zhiyu et al. , 2015
31.6	82.2	~80	15/136	68.4	298.8	2.7	14.2 ± 3.4	Yi Zhiyu et al. , 2015
31.2	84.7	K ₂	33/291	49.0	344.3	5.3	18.2 ± 6.2	Yang Tianshui et al. , 2015
30.9	85.1	99~93	10/82	63.1	224.6	5.1	10.6 ± 6.7	Tang Xiangde et al. , 2013
31.8	87.2	K ₂	8/59	71.2	241.9	5.5/10.0	14.9 ± 8.8	Cao Yong et al. , 2017
32.2	80.8	K ₂	38/274	68.0	211.1	1.9	17.9 ± 2.2	Ma Yiming et al. , 2018
32.2	80.4	116~113	19/205	69.1	319.8	4.8	19.4 ± 5.5	Bian Weiwei et al. , 2017
31.1	81.4	121~117	12/116	70.5	292.9	7.4	15.1 ± 9.1	Yang Tianshui et al. , 2015
32.3	82.6	132~120	51/444	61.4	192.9	2.1	19.9 ± 2.4	Ma Yiming et al. , 2014
31.3	85.1	130~110	18/162	58.2	341.9	4.6	21.3 ± 5.1	Chen Weiwei et al. , 2012

注:SLat. 和 SLon. 分别为采样点的纬度和经度;N/n 为采点数/样品数;Plat. 和 Plon. 分别为古地磁极的经度和纬度;A₉₅为为古地磁极的 95%置信圆锥半顶角;dp/dm 为 95% 置信椭圆半长轴与半短轴;古纬度是以改则参考点(32.5°N / 84°E)计算。

前人根据地质资料对羌塘地块和拉萨地块碰撞和班公湖—怒江洋闭合时限的问题进行了深入探讨,虽然有些学者认为羌塘地块和拉萨地块碰撞及班公湖—怒江洋闭合的时限为中侏罗—早白垩世早期(Xu Ronghua et al. , 1985; Dewey et al. , 1988; Zhu Dicheng et al. , 2016; Huang Tongtong et al. , 2017; Li Shun et al. , 2017; Ma Anlin et al. , 2017, 2018; Ma Yiming et al. , 2018; Sun Gaoyuan et al. , 2019)。但是本文通过对羌塘地块和拉萨地块西部白垩纪古地磁数据的综合分析,显示两地块西部在~115 Ma 时还存在 7.2° ± 5.5°的古纬度差,表明班公湖—怒江特提斯洋西部在早白垩世早期并未闭合。而早白垩世晚期 110~100 Ma 时期羌塘地块和拉萨地块西部的古纬度在误差范围内一致,且晚白垩世时期两地块西部的古纬度已经重叠。以上证据表明羌塘地块和拉萨地块西部发生碰撞以及班公湖—怒江特提斯洋的最终闭合时间在约 110~100 Ma。

班公湖—怒江缝合带西段最高海相地层、蛇绿岩和洋岛等地质证据也支持这一观点。首先,羌塘地块西部存在早白垩世早期的海相地层(Bi Zhiwei et al. , 2018; Chen Weiyan et al. , 2018; Meng Jun et al. , 2018)。Bi Zhiwei et al. (2018)在羌塘地块西南缘发现一套晚侏罗世—早白垩世的海相地层,并获得了该套地层中发育的玄武岩夹层的年龄为 118.3 ± 2.1 Ma。最高海相地层证据表明班公湖—怒江特提斯洋西段在早白垩世早期尚未闭合。其次,班公湖—怒江缝合带西段分布有大量早白垩世蛇绿岩。班公湖—怒江缝合带西段中仓蛇绿岩(114.3 ± 1.4 Ma; Xu Mengjing et al. , 2014)、康穷蛇绿岩(114.6 ± 3.7; Xu Wei et al. , 2015)、洞

措蛇绿岩(~132 ± 3 Ma; Bao Peisheng et al. , 2007)和革吉蛇绿岩(126.6 ± 10.2 Ma; Zhang Kaijun et al. , 2014)的年代均为早白垩世早期,揭示班公湖—怒江缝合带西段在早白垩世早期尚处于洋盆的形成和发展阶段。另外, Baxter et al. (2009)根据班公湖—怒江缝合带西段拉果措蛇绿岩带中放射虫化石的发现揭示青藏高原中部的深海环境可能一直持续到早白垩世 131~121 Ma,并表明班公湖—怒江洋盆闭合发生在早白垩世晚期。最后,班公湖—怒江缝合带西部分布有早白垩世中期洋岛岩石组合(Fan Jianjun et al. , 2014, 2015)。Fan Jianjun et al. (2014, 2015)对班公湖—怒江缝合带西部改则地区仲岗洋岛岩石的地球化学和地质年代学研究表明其形成于以洋壳为基底的大洋洋岛环境,形成时代为早侏罗至早白垩世早期(185~116 Ma)。这表明班公湖—怒江特提斯洋在早白垩世早期仍未闭合(Fan Jianjun et al. , 2018)。

造山带内碳酸盐岩易于受到后期重磁化作用的影响(Huang Wentao et al. , 2017a, 2017b, 2019),而其重磁化与块体间相互碰撞作用有关(Huang Wentao et al. , 2019)。块体的碰撞的强烈作用可能会导致岩石中黄铁矿和磁铁矿氧化生成赤铁矿,从而产生重磁化作用(Huang Wentao et al. , 2019)。羌塘地块与拉萨地块碰撞造成了强烈的构造变形和地壳短缩(Murphy et al. , 1997; Kapp et al. , 2005; Volkmer et al. , 2007)。根据前文分析显示羌塘地块与拉萨地块西部碰撞发生在早白垩世晚期,与本次研究获得的羌塘地块西部晚三叠世灰岩的重磁化时代非常一致。因此,我们认为羌塘地块西部晚三叠世灰岩的重磁化作用与羌塘/拉萨地块的碰撞密切相关。羌塘地块西部晚三叠世灰岩在

早白垩世晚期发生重磁化作用也是羌塘地块与拉萨地块西部碰撞的响应。

4 结论

(1)本次研究从羌塘地块西部改则地区晚三叠世灰岩中获得的其特征剩磁分量未通过褶皱检验,表明为后期重磁化的结果。研究剖面高温特征剩磁平均方向在地理坐标下平均方向为 $D_g = 349.3^\circ$, $I_g = 40.4^\circ$, $\kappa_g = 45.4$, $\alpha_{95} = 6.5^\circ$, 相应的古地磁极为 76.4°N , 311.1°E , $dp/dm = 4.7^\circ/7.9^\circ$ 。这一古地磁极与羌塘地块早白垩世约 110~100 Ma 的古地磁极重叠,由此确认本次研究地层重磁化的时代为早白垩世 110~100 Ma。

(2)结合羌塘地块现有可靠的古地磁结果表明羌塘地块西部南缘在约 110~100 Ma 期间位于 $23.3 \pm 4.0^\circ\text{N}$ 的古纬度位置。对比拉萨地块西部白垩纪古地磁数据并结合海相地层、蛇绿岩和洋岛等地质证据,提出羌塘地块和拉萨地块西部碰撞以及班公湖—怒江特提斯洋西段闭合的时间在早白垩世晚期约 110~100 Ma。

(3)羌塘地块西部改则地区晚三叠世灰岩的早白垩世晚期重磁化作用与羌塘/拉萨地块西部碰撞密切相关。

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Paleomagnetic study of the Late Triassic limestones in the western portion of the Qiangtang Block and its tectonic implications

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Abstract

The processes of the collision between the Qiangtang and Lhasa blocks is a key factor to understanding the evolution history and geodynamics of the central Tibetan Plateau. The paleomagnetic study for the Qiangtang block is critical to understand these questions. A paleomagnetic study of the Late Triassic limestones from the Gaize area in the western portion of the Qiangtang block has been carried out and a remagnetization component has been obtained. The mean direction of the characteristic remanent component in geographic coordinates is $Dg = 349.3^\circ$, $Ig = 40.4^\circ$, $\kappa g = 45.4$, $\alpha_{95} = 6.5^\circ$, corresponding to a paleopole at 76.4°N , 311.1°E with $dp/dm = 4.7^\circ/7.9^\circ$. This pole coincides with the Early Cretaceous (ca. 110~100 Ma) paleomagnetic pole of the Qiangtang block, indicating the age of the remagnetization is 110~100 Ma. Our new results, combined with previous reliable Cretaceous paleomagnetic results from the Qiangtang and Lhasa blocks and synthesize stratigraphy, ophiolite and ocean island evidence, suggested that the final closure of western portion of the Bangong-Nujiang Tethys Ocean may have occurred during 110~100 Ma. Furthermore, the late Early Cretaceous remagnetization of the Late Triassic limestones in the Gaize area was closely related with the Qiangtang-Lhasa collision.

Key words: Qiangtangblock; paleomagnetism; remagnetization; Qiangtang-Lhasa collision; Bangong-Nujiang Ocean